

埃达克质岩的金属成矿作用

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摘要:介绍了“埃达克质岩”的术语、与成矿有关的埃达克质岩的分布、成矿背景,讨论了埃达克质岩有利于成矿的控制因素。“埃达克质岩”是指那些具有与俯冲洋壳熔融形成的“埃达克岩”类似地球化学特征,如 $\text{SiO}_2 \geq 56\%$, $\text{Al}_2\text{O}_3 \geq 15\%$, 亏损Y ($\leq 18 \times 10^{-6}$)和重稀土元素(如 $\text{Yb} \leq 1.9 \times 10^{-6}$),高Sr(很少样品的Sr含量低于 400×10^{-6}),无-正Eu, Sr异常,贫高场强元素等,但可以形成于不同构造背景并可有不同成因的岩浆岩。埃达克质岩具有重要的金属成矿意义,其有利成矿背景主要包括岛弧、大陆板内伸展和大陆活动碰撞造山带环境。世界上许多(包括三个最大的)斑岩铜矿都与埃达克质斑岩密切共生,因此埃达克质岩的成矿潜力巨大。在岛弧和大陆板内伸展环境中,来自俯冲玄武质洋壳或洋壳沉积物或拆沉的大陆地壳产生的熔体或释放的超临界流体与地幔的相互作用,一方面可能导致熔体被地幔橄榄岩混染,另一方面可能导致高 Fe_2O_3 含量的熔体或超临界流体对地幔的交代作用,地幔氧逸度升高,地幔金属硫化物被氧化分解,有利于铜、金等的矿化。

关键词:埃达克质岩;铜金矿化;斑岩铜矿;地球动力学

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The Metalliferous Mineralization Associated with Adakitic Rocks

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Abstract: Adakitic rocks have important implications for the metalliferous mineralization. This paper introduces the nomenclature of “adakitic rocks”, the distribution and tectonic settings of the adakitic rocks associated with metalliferous mineralization. The factors that the adakitic rocks are favorable for the metalliferous mineralization are discussed in detail at last. Similar to a subducted oceanic crust-derived adakite, the “adakitic rocks” are characterized by the following geochemical characteristics: $\text{SiO}_2 \geq 56$ wt%, $\text{Al}_2\text{O}_3 \geq 15$ wt%, depleted Y ($\leq 18 \times 10^{-6}$) and heavy rare earth element contents ($\text{Yb} \leq 1.9 \times 10^{-6}$), high Sr (rarely $< 400 \times 10^{-6}$), negligibly positive Eu and Sr anomalies, and depleted high field strength element contents. They can be generated in different tectonic settings and by different mechanisms. The favorable tectonic settings for the metalliferous mineralization include arc, within-continent extension, and active collisional orogenic belt between different continents. Many large and giant are closely associated with adakitic rocks, indicating that they have a huge metalliferous potential. In the arc and within-continent extension settings, the interaction between subducted basaltic oceanic crust, or sediments, or delaminated continental crust-derived melts, or released super critical fluids with mantle causes contamination of melts by mantle peridotites on the one side, and metasomatism of mantle peridotites by the melts or fluids with high Fe_2O_3 on the other side. Thus, the raised oxygen fugacity in the mantle causes metal sulfides to be decomposed, which is favorable for the Cu-Au mineralization.

Key words: adakitic rocks; Cu-Au mineralization; porphyry copper deposits; geodynamics

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1 引言

埃达克岩(adakite)是Defant & Drummond (1990)在《Nature》上发文提出的一类由俯冲的年轻且热的洋壳在榴辉岩相条件下熔融形成的、具有特殊地球化学特征的中酸性火成岩。该概念的提出引起了广泛关注,也引起了一定争议,主要原因是埃达克岩与其它成因的岩石具有类似的地球化学特征。目前提出的这类岩石的成因机制主要包括:(1)俯冲洋壳的熔融(Defant et al, 2002; Li and Li, 2003; Zhang et al, 2006; 李武平, 2006; Zhou et al, 2006a; Zhao et al, 2007; Wang et al, 2007a, 2008a); (2)含水玄武岩发生含石榴石的高压分离结晶(Macpherson et al, 2006); (3)地壳混染-(低压)分离结晶(AFC)(Castillo et al, 1999; Castillo, 2006)或分离结晶(Rodríguez et al, 2007); (4)熔融-混染-存储(storage)-均化(MASH)过程(Richards and Kerrich, 2007); (5)长英质和玄武质岩浆的混合(Wang et al, 2003a; Guo et al, 2007a; Qin et al, 2007; Streck et al, 2007); (6)增厚下地壳的熔融(Atherton and Petford, 1993; 张旗等, 2001; Liu et al, 2003; Xiong et al, 2003; Hou et al, 2004; Chung et al, 2003, 2005; Wang et al, 2005, 2006a, 2007b; Xue et al, 2006; Xu et al, 2007; Zhang et al, 2006b, 2007; Shi et al, 2007; Jiang et al, 2007, Lai et al, 2007; Xiao et al, 2007; Zhao et al, 2008); (7)拆沉下地壳的熔融(Xu et al, 2002, 2008; Gao et al, 2004; Wang et al, 2006b; Hou et al, 2007; Liu et al, 2008; Yang and Li, 2008; Huang et al, 2008); (8)俯冲陆壳的熔融(Wang et al, 2008b); (9)富集地幔的低度熔融(Jiang et al, 2006)或俯冲板片熔体交代的地幔的熔融(Martin et al, 2005; Gao et al, 2007; Li et al, 2008),等等。尽管存在激烈的争论,但由于这种岩石可能与俯冲洋壳或陆壳的熔融、增厚或拆沉下地壳的熔融、玄武质岩浆的底侵、板片窗(slab window)以及地壳生长、MASH、AFC、岩浆混合等过程密切相关,因此具有重要的地球动力学意义。

近年来,具有与埃达克岩类似地球化学特

征的岩石受到关注的另外一个重要原因就是,许多与铜-金-钼-铁等金属矿化与这类岩石密切共生,世界上一些最大的斑岩铜矿床(El Teniente, Chuquicamata, Río Blanco-Los Bronces等)、中国几个最大铜矿床(如驱龙、德兴、玉龙、土屋-延东等)的成矿斑岩都被认为与埃达克岩具有类似的地球化学特征。一些研究者甚至认为,这类岩石可以作为找矿标志,用于铜-金的勘探(Defant and Kepezhinskis, 2001; 张旗等, 2004)。作者近年来开展了这类岩石的一些研究工作,本文将在总结近年来有关这类岩石的成矿作用研究工作的基础上,结合自己收获的资料,来探讨这类岩石影响金属成矿的可能因素。

2 “埃达克质岩”的术语

在讨论成矿作用之前,有必要简单回顾“埃达克岩”的术语。埃达克岩的命名是一个有争议的话题,引起争议的最重要原因是该类岩石的命名并不是按照传统的岩石命名方法,即以岩相学研究为基础的矿物组成来命名。正如Defant & Drummond (1990)所指出的那样,埃达克岩的岩相学特征是有些变化的,其矿物组成是斜长石+角闪石±黑云母±辉石±不透明矿物,具有这样矿物组合的岩石可以包括安山岩-英安岩-富钠流纹岩(ADR)及相应的侵入岩。Defant & Drummond (1990)在命名“埃达克岩”时强调了3个原则:(1)这类由俯冲洋壳熔融形成的岩石最先在阿拉斯加阿留申群岛的Adak岛被发现(Kay, 1978),因此以其地名来命名,实质上包含了岩石成因的含义;(2)这类岩石是与俯冲的年轻(≤ 25 Ma)大洋岩石圈共生的火山岩或侵入岩,强调了其形成于岛弧的构造背景;(3)这类岩石具有如下的地球化学特征: $\text{SiO}_2 \geq 56\%$, $\text{Al}_2\text{O}_3 \geq 15\%$ (很少低于15%), MgO 通常 $< 3\%$ (很少 $> 6\%$),低于正常岛弧ADR的Y($\leq 18 \times 10^{-6}$)和重稀土元素(HREE,如Yb $\leq 1.9 \times 10^{-6}$),高于正常岛弧ADR的Sr(很少 $< 400 \times 10^{-6}$),无-正Eu, Sr异常,同正常岛弧ADR一样贫高场强元素(HFSE),具有低的初始 $^{87}\text{Sr}/^{86}\text{Sr}$ 比值(< 0.704)。

按照Defant和Drummond (1990)的定义,“埃达克岩”的原始含义应该包含了“岩石成

因”、“构造背景”与“地球化学特征”三个方面。如果严格按照上述定义,很少有岩石能够被称为“埃达克岩”,主要表现在:(1)岩石的元素-同位素地球化学特征、构造背景与“埃达克岩”完全类似,但成因存在争论,如菲律宾南部、墨西哥Baja地区的一些新生代中酸性岩浆岩,当它们被认为是俯冲洋壳熔融的时候,它们可以被称为“埃达克岩”(Sajona et al, 1993, 1996; Aguillon-Robles et al, 2001),但是当它们被认为是由其它过程(如AFC过程, Castillo et al, 1999; 增厚地壳熔融, Castillo, 2008)形成时,其名称则是有疑问的;(2)岩石的主、微量元素特征与“埃达克岩”类似,但俯冲板片的年龄、Sr同位素特征乃至可能的成因与“埃达克岩”的不一致,如著名的秘鲁Codillera Blanca杂岩具有与“埃达克岩”类似的主量、微量元素特征,但是与其共生的俯冲洋壳大于60 Ma且岩石的初始 $^{87}\text{Sr}/^{86}\text{Sr}$ 比值 >0.705 ,当有些研究者认为其为岛弧环境中新底侵的玄武质洋壳熔融形成时,其不能被称为“埃达克岩”而是被称为“富Na的熔体”(Atherton and Petford, 1993; Petford and Atherton, 1996),但当有研究者认为其可以由平坦(flat)俯冲的老洋壳熔融形成的时候,又可以被称为“埃达克岩”(Gutscher et al, 2000);(3)一些中酸性岩浆岩,既不与俯冲年轻洋壳共生(或形成于岛弧环境),且同位素地球化学特征以及某些主量元素(如 K_2O)也与“埃达克岩”差别较远,仅仅一些重要的主、微量元素特征(如 SiO_2 , Al_2O_3 , MgO , Y, HREE, Sr和HFSE含量和Eu, Sr异常)与“埃达克岩”类似,因此其定名争议则更大。但是,现实的情况是,第三类岩石在自然界中出露则更为广泛(张旗等, 2001, 2003; Castillo, 2006)。近年来,许多研究者把不能严格符合“埃达克岩”的原始含义的岩石用一些新的名称来代替,如高Ba-Sr花岗岩类(Tarney and Jones, 1994; Qian et al, 2003)、埃达克质(adakitic; Stern et al, 1996; Yunul et al, 2000)或类埃达克(adakite-like; Harris et al, 1996)岩、“C型埃达克岩”(张旗等, 2001)或“地壳起源的(Crustally-derived)埃达克岩”(Sheppard et al, 2001)、富Sr安山岩(Conrey et al, 2001)以及高

Sr/低Y型中酸性火成岩(Ge et al, 2002)等等。这些名称表明,上述研究者已经充分注意到所描述的岩石与“埃达克岩”的原始含义的差别。

作者认为,作为一个岩石名称不应该以岩石的成因、构造背景为前提,而应该反映该岩石最基本的特征,这个特征不会随人们对岩石成因、构造背景的认识不同而改变(Wang et al, 2007b)。经典的岩石名称都是以岩相学特征为基础的,但是仅仅依靠岩相学往往使一些具有特殊地球化学特征的岩石很难被识别出来。随着近30~40年来地球化学测试手段的突飞猛进,一些具有特殊地球化学特征的岩石相继被识别出来,如A型花岗岩(Loiselle and Wones, 1979)、巴哈岩(bajaites; Rogers et al, 1985)、高 $\text{Mg}^\#$ 安山岩(Kelemen, 1995)、高铝玄武岩(Hagerty et al, 2005)、富Nb玄武岩(Sajona et al, 1993)等,一些古老的岩石名称又重新被关注,如“玻安岩”(boninite; Cameron et al, 1979)、赞岐岩(sanukite)或赞岐岩类sanukitoid(Shirey and Hanson, 1984; Tatsumi and K, 1981)等。上述岩石类型基本上需要地球化学资料才能确定其岩石名称,它们也与“埃达克岩”一样,都具有极其重要的地球动力学意义。

作者认为,“埃达克岩”的原始概念中最核心的特征,应该是这些地球化学特征: $\text{SiO}_2 \geq 56\%$, $\text{Al}_2\text{O}_3 \geq 15\%$, 亏损Y($\leq 18 \times 10^{-6}$)和HREE(如, $\text{Yb} \leq 1.9 \times 10^{-6}$), 高的Sr(很少样品 $< 400 \times 10^{-6}$), 无正Eu, Sr异常, 贫HFSE。因此,作者把具有上述地球化学特征的岩浆岩称为“埃达克质(adakitic)岩”,而不考虑这类岩石成因、构造背景以及其它的地球化学特征(如 K_2O , MgO 含量和同位素比值等)(Stern et al, 1996, Yunul et al, 2000; Hou et al, 2004; Ishihara et al, 2005; Wang et al, 2005, 2006a, b, c, 2007a, b; Rodriguez et al, 2007; Zhao and Zhou, 2007; Xu et al, 2008; Hou et al, 2007; Huang et al, 2008; Lai et al, 2007; Yang and Li, 2008)。也就是说,“埃达克质岩”的含义更广,包含了所有具有上述地球化学特征但可能形成不同构造背景并有不同成因的岩浆岩,甚至包括了经典的“埃达克岩”(Defant and Drummond, 1990)。

3 与成矿有关的埃达克质岩的分布与成矿背景

Thiélemont等(1997)最先总结并报道了世界上的许多Cu-Au-Mo-Ag矿床与埃达克质岩密切共生。Sajona and Maury(1998)则报道了菲律宾绝大多数Cu-Au-Mo-Ag矿床都与埃达克质岩以及相关的岩石(如富Nb玄武岩)共生。此后的2001~2003年,包括中国学者在内许多研究者都发现了埃达克质岩与Cu-Au-Mo-Ag矿化密切共生的现象(Defant and Kepezhinskias, 2001; Oyarzun et al, 2001; Defant et al, 2002; Mungall, 2002; 王强等, 2001a, b, 2003; 张旗等, 2001, 2003; 侯增谦等, 2003; Bissig et al, 2003; Morozumi, 2003; Reich et al, 2003, Gonzalez-Partida et al, 2003; Wang et al, 2003; Yumul et al, 2003; Bissig et al, 2003; Defant et al, 2002; Gonzalez-Partida et al, 2003; Morozumi, 2003; Mungall, 2002; Oyarzun et al, 2001; Reich et al, 2003; Wang et al, 2003b; Yumul et al, 2003)。特别是,自2003年以后,我国学者陆续报道了许多与Cu-Au-Mo-Ag-Fe矿化密切共生的埃达克质岩(参考文献转引自王强等,

2007)。Cu-Au-Mo-Ag-Fe矿化与埃达克质岩的密切共生关系使得一些学者提出,埃达克质岩可以用于铜金勘探(Defant and Kepezhinskias, 2001; 张旗等, 2004),就象在金伯利岩中找金刚石(Defant and Kepezhinskias, 2001)。

3.1 与埃达克质岩有关的金属矿床的分布

与埃达克质岩有关的金属矿床在全球广泛分布,尤其以环太平洋地区分布最广(图1),且很多都是大型-超大型-最大型矿床(表1)。世界上最大的三个斑岩Cu(-Mo-Au)矿床——智利的El Teniente, Chuquibambilla, Rfo Blanco-Los Bronces矿床的铜储量分别达到 94×10^6 , 66×10^6 和 57×10^6 吨,都位于环太平洋岛弧带,都与新生代俯冲洋壳熔融形成的埃达克质斑岩密切共生(Thiélemont et al, 1997; Cooke et al, 2005)。我国与埃达克质斑岩密切共生的Cu-Mo-Au-Fe矿床主要位于三大成矿域:环太平洋的中国东部、青藏高原及周边、中亚造山带(图2)。同世界上的一些超大型-最大型矿床相比(表1),中国的斑岩铜矿床则小得多(表2),如中国最大的四个斑岩Cu(-Mo-Au)矿床——德兴、土屋-延东、驱龙与玉龙的铜储量分别为 8.3×10^6 , 7.74×10^6 , 7.05×10^6 和 6.22

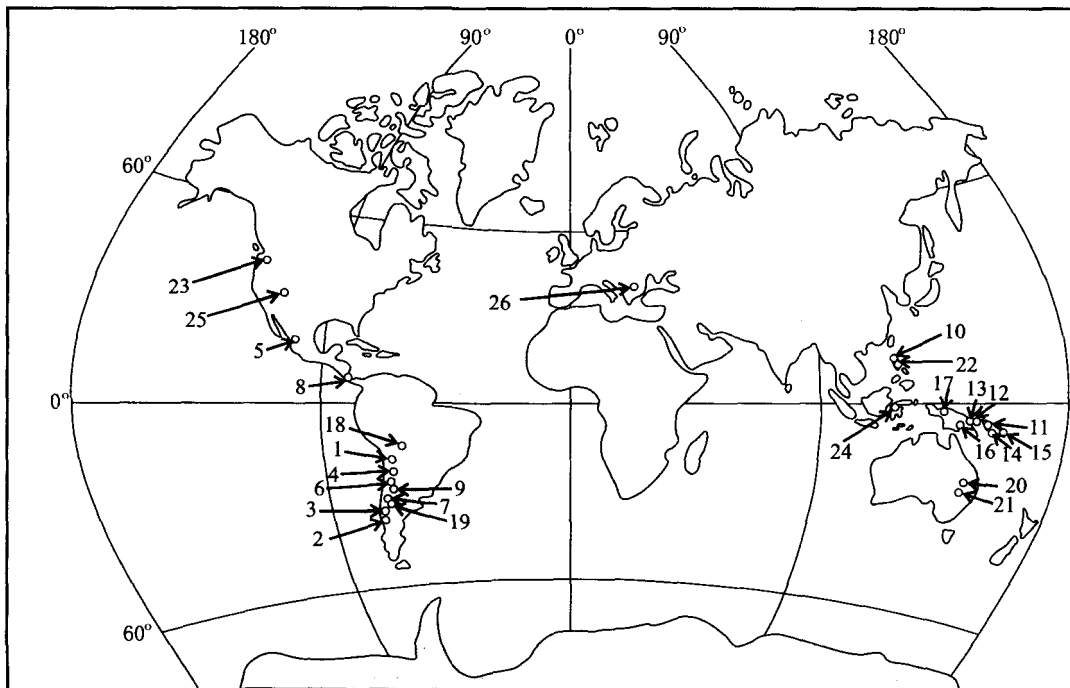


图1 世界上与埃达克质岩有关的Cu-Au-Mo-Fe矿床分布(详细名称及文献见表1)

Fig. 1 The distribution of Cu-Au-Mo-Fe deposits associated with adakitic rocks in the world

$\times 10^6$ 吨,但都与埃达克质斑岩密切共生(如, Hou et al, 2004; Jiang et al, 2006; Wang et al, 2003, 2006; Zhang et al, 2006)。近年来,在我国新疆新发现的土屋-延东,藏南的甲马、厅宫、冲江等,以及云南的普郎铜矿等都与埃达克质岩密切共生。尽管这些与金属矿化密切相关的埃达克质岩的成因还存在激烈的争论(Richards and Kerrich, 2007),但是Cu-Mo-Au-Fe矿化与埃达克质岩共生的现象、埃达克质岩成矿潜力值得关注。

3.2 成矿背景

3.2.1 岛弧背景埃达克质岩的成矿

几乎所有环太平洋地区新生代的斑岩Cu(-Mo-Au)矿床都产出在岛弧环境,且那些与埃达克质岩共生金属矿床的形成主要可能与俯冲板片的熔融有关。但是,对于板片发生熔融的动力学模式则有明显的差别,概括起来,主要有:

(1)洋脊俯冲与板片窗;(2)海山链、大洋高原的平坦俯冲;(3)年轻洋壳的俯冲等。我国新疆地区的许多晚古生代与Cu(-Mo-Au)矿床密切共生的埃达克质岩也主要产出在岛弧环境(Zhang et al, 2006; Han et al, 2006; 王强等, 2006; Wang et al, 2007)。

3.2.2 大陆板内伸展背景

国际上报道的产于陆内伸展背景中且与埃达克质岩密切共生的Cu-Mo-Au矿床并不是很多(Ishihara et al, 2005; Pinto-Linares et al, 2008)。近年来开展埃达克质岩与金属矿化关系研究的一个典型地区是扬子地块东部。该区存在大量燕山期埃达克质岩与Cu-Mo-Au-Fe矿床密切共生。尽管该区燕山期的构造背景存在争议(Xu et al, 1999; Li, 2000; Zhou and Li, 2000; Zhou et al, 2006; Sun et al, 2007; Li and Li, 2007),但根据区内出现燕山期变质核杂岩、A型或伸展型岩浆岩以及断陷盆地的出现,作者趋向认为扬子地块东部燕山期可能处于一个板内伸展的背景中(Li, 2000),并且区内断裂带在这些岩浆岩的形成中可能发挥了重要作用(周新民等, 2006)。这样,认为扬子地块东部燕山期埃达克质岩及共生的Cu-Mo-Au-Fe矿床可能形成于板内伸展的背景中。

3.2.3 大陆活动碰撞造山带

青藏高原是世界上最大、最高的高原,也是陆-陆碰撞形成的最典型的碰撞造山带。印度-欧亚大陆的碰撞自~70~40 Ma开始,大量的GPS资料显示两大陆间的汇聚现今仍旧在进行,也表明青藏高原为一典型的构造造山带。在该活动造山带中主要发育了两期埃达克质岩:始新世(46~36 Ma; Jiang et al, 2006; Liu et al, 2003; Wang et al, 2008)和渐新世—中新世(26~9 Ma; Chung et al, 2003; Hou et al, 2004; Qu et al, 2004; Gao et al, 2007; Guo et al, 2007b; Mo et al, 2007; Lai et al, 2007)。与这些岩石对应,青藏高原Cu-Mo-Au矿床形成时代也主要在这两个时期(Hou et al, 2006; 侯增谦等, 2005),并且大多数与埃达克质岩共生。

4 讨论

埃达克质岩与金属成矿关系密切,但也并不是所有的埃达克质岩都与金属矿化密切共生(Wang et al, 2007c)。许多研究者从不同的角度讨论了埃达克质岩有利于成矿的可能控制因素。张旗等(2004)认为制约埃达克(质)岩成矿的关键因素在于:在角闪岩相向榴辉岩相转变过程中发生脱水产生埃达克质岩浆,同时萃取地幔中的成矿元素进入岩浆。刘红涛等(2004)认为埃达克(质)岩成矿的有利因素在于埃达克质岩浆的富流体、高氧逸度和基性源岩等特征。侯增谦等(2005)认为藏南冈底斯中新世由于来自深部亏损地幔物质与下地壳物质交换作用,不仅导致冈底斯地壳加厚、下地壳熔融形成埃达克质斑岩,而且提供了巨量金属供应。但有学者则认为冈底斯斑岩铜矿与雅鲁藏布俯冲板片及其上的少量洋壳沉积物混合物质部分熔融有关(Qu et al, 2004)或与板片熔体交代的上地幔熔融有关(Gao et al, 2007)。Zhang et al(2006)和Han et al(2006)则强调了俯冲板片熔融对成矿的贡献。

前面总结成矿的埃达克质岩的构造背景发现,这种岩石可以形成三种构造背景中:岛弧、大陆板内伸展和大陆活动造山带。这表明埃达克质岩的成矿并不完全受构造背景制约,而更可能与岩石的特殊性有关。另外,作为亲硫元素,

表2 国内与埃达克质岩有关的Cu-Au-Mo-Fe矿床
Table 2 Cu-Au-Mo-Fe deposits associated with adakitic rocks in China

序号	矿床	地点	Cu (Mt)	Mo (Mt)	Fe (Mt)	Au (t)	Ag (t)	时代 (Ma)	文献
1	土屋-延东	东天山	7.74			0.187		343~323	(1)(2)
2	包古图	西准	0.6			200		310	(3)
3	三岔口	东天山	0.7					278	(4)
4	阿吾拉勒	西天山	1.75					247~268	(5)
5	卡拉先格尔	富蕴南	1.67					381	(6)
6	冲江	西藏尼木	0.67	0.018		20.09	397.1	12.9	(7)(8)
7	厅官	西藏尼木	0.3					13.5	(9)(10)
8	南木	西藏曲水						14.67	(7)(10)
9	拉抗俄	西藏达孜						13.5~13.6	(7)(10)
10	甲马	西藏墨竹工卡						15.18	(7)(11)
11	驱龙	西藏墨竹工卡	7.05	0.455			5340	15.82~16.85	(7)(12~13)
12	马拉松多	西藏察雅	1					36~35.4	(7)(14)
13	扎位尕	西藏察雅	0.3					34~40	(7)(14)
14	莽总	西藏察雅	0.25					34~41	(7)(14)
15	多霞松多	西藏察雅	0.5					41~52	(7)(14)
16	玉龙	西藏察雅	6.22						(7)(14)
17	普朗	云南香格里拉	1.14	0.01		0.03		228	(15)
18	雪鸡坪	云南香格里拉	0.19			3		249.92	(16~17)
19	西范坪	云南香格里拉	0.18					32.1	(17)
20	马场箐	云南武定	0.25					35.8	(17)
21	德兴	赣东北	8.3	0.23		138	773	171	(18)
22	大冶铁山	鄂东南	0.70		160			134	(19)
23	封山洞	鄂东南						130	(19)
24	铜山口	鄂东南	0.42					147	(19)
25	滁州	安徽	0.17			8.1		127	(20)
26	安基山	江苏	0.165					123~106	(21)
27	月山	安徽						136	(19)(22)
28	沙溪	安徽	0.32					136	(23)
29	总铺	安徽						132	(19)
30	利国	江苏						131	(23)
31	多宝山	东北	2.97	0.11				485	(24)
32	团结沟	东北				52.7		105	(24)
33	弓棚子	东北	0.20					中生代	(24)
34	大黑山	东北		1.09	65			170	(24)(26)
35	小西南岔	东北	0.11			31		134~130	(24)
36	白乃庙	东北	0.51	0.8		19.2	181.7	440~386	(24)
37	涑源	华北	0.09			27.5	1224	142	(24)
38	玲珑	胶东				24.2		119	(24~25)

(1) Zhang et al, 2006; (2) Rui et al, 2002; (3) 宋会侠等, 2007; (4) 李华芹等, 2004; (5) Zhao et al, 2008; (6) 万博等, 2006; (7) 侯增谦等, 2003a; (8) 林武等, 2004; (9) 李金祥等, 2007; (10) 侯增谦等, 2003b; (11) 李光明等, 2005; (12) 孟祥金等, 2003; (13) 王亮亮和莫宣学, 2006; (14) Hou et al, 2003; (15) 冷成彪等, 2007; (16) 曾普胜等, 2003; (17) 曾普胜等, 2006; (18) Wang et al, 2006b; (19) Wang et al, 2007c; (20) 资锋等, 2007; (21) Xu et al, 2002; (22) Sun et al, 2003; (23) Wang et al, 2006c; (24) Xu et al, 2004; (25) 张旗等, 2003; (26) 李俊建等, 2005; (27) Ge et al, 2007.

铜、金等主要以金属硫化物的形式存在于地幔中 (Mungall, 2002), 这些幔源的金属物质如何从地幔上升到地壳中并成矿, 也是我们必须考虑的重要方面。一般认为, 金属硫化物只有在高氧逸

度的情况下才能发生分解进入到流体或岩浆中, 然后到达地表成矿 (Mungall, 2002)。作者以天山北部石炭纪岛弧埃达克岩、扬子地块东部燕山期埃达克质岩和藏南冈底斯中新世埃达克质岩分

别代表形成于上述三种背景中的岩石, 来初步探讨埃达克质岩为什么有利金属成矿的原因。

图3显示, 来自天山北部石炭纪岛弧成矿埃达克岩、扬子地块东部燕山期成矿埃达克质岩都显示了高MgO, $Mg^{\#}$ 和低 FeO^*/MgO 比值(图3a-f), 但扬子地块东部燕山期贫矿埃达克质岩却显示了低的MgO, $Mg^{\#}$ 和高的 FeO^*/MgO 比值(图3d-f)。这些特征表明, 上述两个地区的成矿埃达克质岩可能包含有较多的地幔组分, 它们的成因很可能与“熔体-地幔作用”有关(王强等, 2006; Wang et al, 2006a, b, 2007c)。

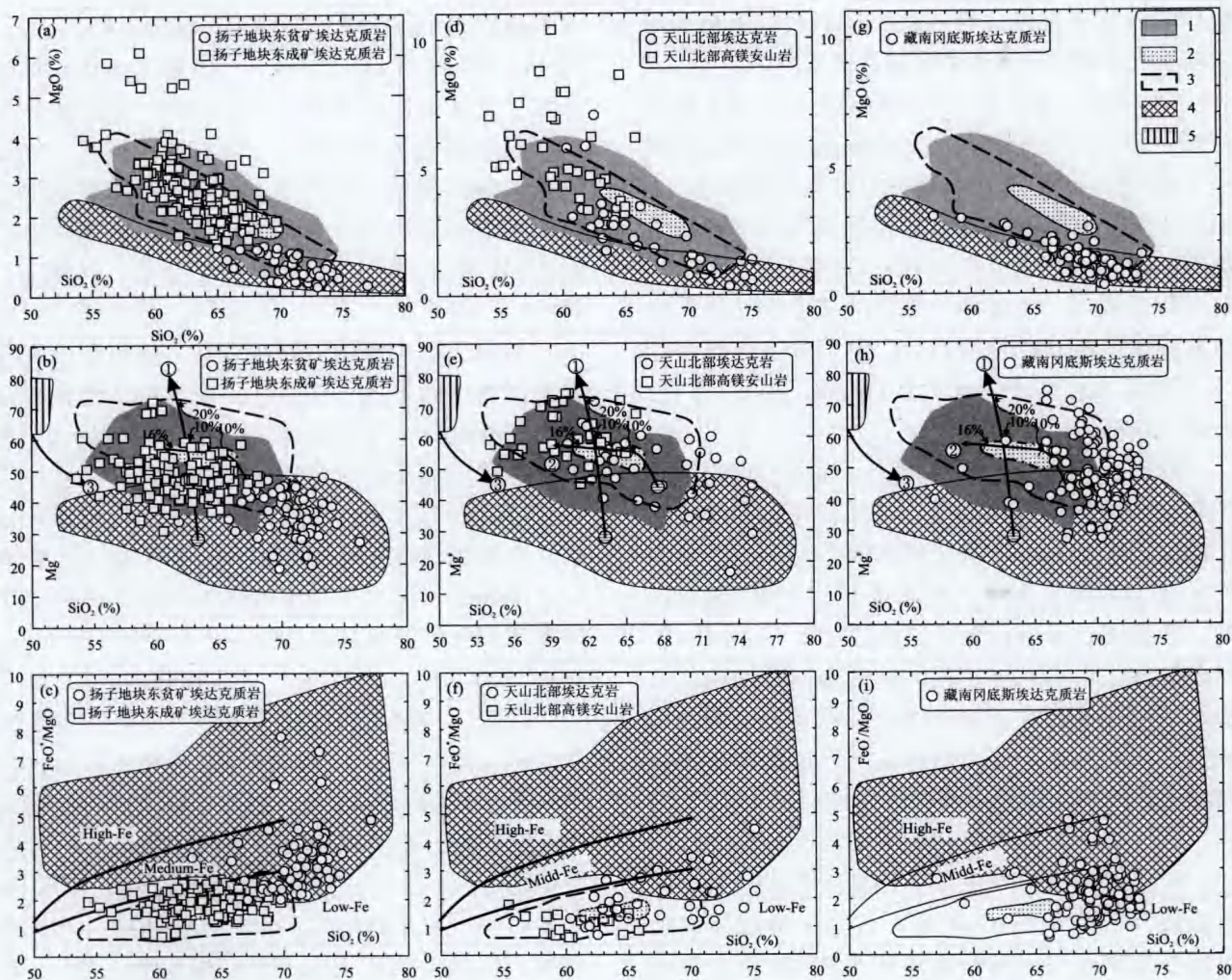
在岛弧背景中, 俯冲洋壳(或沉积物)熔融产生的熔体与地幔楔橄榄岩之间存在强烈的相互作用。以天山北部石炭纪的岛弧为例, 一些重要的矿床如喇嘛苏铜矿、阿希金矿、土屋-延东铜矿、赤湖铜矿等与石炭纪的埃达克质(adakitic)岩及其共生岩石共生, 我们认为向南俯冲的北天山洋的年轻洋壳在榴辉岩相的条件下发生熔融形成埃达克质熔体, 这类熔体在上升过程中与地幔楔橄榄岩的相互作用形成埃达克岩-高镁安山(闪长)岩(如赞岐岩类)-富Nb玄武质岩组合, 同时, 由于俯冲板片产生的埃达克质岩浆具有高的氧逸度(即高的 Fe_2O_3 含量), 其与地幔楔橄榄岩的强烈相互作用将导致储存于地幔中的金属硫化物发生氧化分解, 释放出亲铜元素(如铜、金等)进入到岩浆中(图4a)。这是天山北部铜金矿床(如喇嘛苏、阿希、土屋-延东、赤湖, 等等)与一些埃达克岩、高镁安山(闪长)岩或富Nb岛弧玄武质岩密切共生的一个重要原因(王强等, 2006)。近期的大量研究也表明, 俯冲板片在高温高压下熔融形成熔体或释放的超临界流体含有丰富的板片组分(如大离子亲石元素、轻稀土元素、Th, U)以及 Fe_2O_3 (Mungall, 2002; Kepezhinskias et al, 2002; Widom et al, 2003; Borisova et al, 2006; Portnyagin et al, 2007)。这使得产生的岩浆或流体具有比地幔楔橄榄岩明显高的氧逸度, 这些岩浆或流体对地幔楔橄榄岩的交代作用将明显提高地幔的氧逸度(图4b)(Mungall, 2002; Bryant et al, 2007)。

在板内伸展背景中, 拆沉榴辉岩质下地壳或残留洋壳熔融产生的熔体与地幔楔橄榄岩之间也

存在强烈的相互作用(Xu et al, 2002; Gao et al, 2004; Wang et al, 2006)。以扬子地块东部的长江中下游地区以及德兴地区为例, 该区被誉为“我国东部的工业走廊”, 其在燕山期为一典型的陆内伸展背景。扬子地区一些铜金成矿作用与燕山期埃达克质岩密切共生(如铁山、铜碌山、铜山口、封山洞、沙溪、月山、总铺、铜官山、狮子山、冬瓜山、铜厂、富家坞、朱砂红、安基山、滁州、冶山等), 这些岩石由拆沉下地壳或残留洋壳熔融形成。作者认为由拆沉下地壳或残留洋壳熔融产生具有高的氧逸度(即高的 Fe_2O_3 含量)的埃达克质岩浆与地幔楔橄榄岩的相互作用也将导致储存于地幔中的金属硫化物发生氧化分解, 释放的亲铜元素(如铜、金等)进入到岩浆中(图4a)。这也是扬子地块东部许多燕山期埃达克质岩有利于铜金成矿的一个重要原因(Wang et al, 2007c)。

因此, “熔体-地幔作用”具有非常重要的动力学与成矿意义, 在熔体-地幔相互作用过程中: (1) 熔体的多来源, 可以来自玄武质洋壳或洋壳沉积物, 也可以来自拆沉的大陆地壳, 其成分是复杂的; (2) 与“熔体-地幔作用”相联系的动力学过程是复杂的, 包括洋壳俯冲、下地壳拆沉、地幔上涌或地幔交代或熔融等; (3) 熔体被地幔楔橄榄岩混染, MgO增高, 但高 Fe_2O_3 含量的熔体对地幔交代作用在地幔中形成角闪石-钛铁矿或角闪石-富Fe辉石, 熔体的 FeO/MgO 可能降低, 但地幔氧逸度升高, 这将导致地幔金属硫化物分解, 有利于铜金矿化。

然而, 对位于大陆活动造山带中的冈底斯成矿埃达克质斑岩, 情况可能比较复杂。在图3中, 它们总体显示了较低MgO含量, 但部分样品显示的高 $Mg^{\#}$ 和 FeO^*/MgO 值, 其成因可能与增厚下地壳熔融(Chung et al, 2003; Guo et al, 2007b)、幔源与下地壳源物质的交换(Hou et al, 2004; 侯增谦等, 2005)、俯冲板片及其上的少量洋壳沉积物混合物部分熔融(Qu et al, 2004)或与板片熔体交代的上地幔熔融(Gao et al, 2007)有关。很显然, 冈底斯成矿埃达克质斑岩的成因仍旧存在激烈的争论, 这个争论也制约着对该区成矿作用的进一步深入探讨。



图中一些典型埃达克质岩、实验熔体的成分范围转引自Wang et al (2006b, 2007c): 1-折沉下地壳产生的埃达克质岩, 2-受橄辉岩混染的变玄武岩和榴辉岩的实验熔体, 3-俯冲洋壳熔融形成的埃达克岩, 4-变玄武岩和榴辉岩的实验熔体(1-4.0 GPa), 5-地幔熔体; ①地幔AFC曲线, ②地幔AFC曲线, ③地壳AFC曲线; 扬子地块东部燕山期埃达克质岩的样品引自Wang et al (2006b, c; 2007b, c); 天山北部地区石炭纪埃达克质岩、高镁安山岩的样品主要引自王强等(2006), Wang et al (2007a), Zhang et al (2006a)及所引参考文献; 藏南冈底斯渐新世-中新世埃达克质岩的样品引自Chung et al (2003), Hou et al (2004), Qu et al (2004), Gao et al (2007)和Guo et al (2007b).

The composition fields of typical adakitic rocks and experimental melts are from Wang et al. (2006b, 2007c): 1-Delaminated lower crust-derived adakitic rocks; 2-metabasalt and eclogite melts contaminated by peridotites; 3-subducted oceanic crust-derived adakitic rocks; 4-metabasalt and eclogite melts (1-4.0 GPa); 5-mantle melts; ① mantle AFC curve; ② mantle AFC curve; ③ crustal AFC curve. The Yanshanian adakitic rock samples in the eastern Yangtze block are from Wang et al (2006b, c, 2007b, c); The Carboniferous adakitic rock and high-Mg andesite samples in the northern Tianshan area are from Wang et al (2006, 2007a), Zhang et al (2006a) and references therein; the Oligocene-Miocene adakitic rock samples in the Gangdese area (South Tibet) are from Chung et al (2003), Hou et al (2004), Qu et al (2004), Gao et al (2007) and Gou et al (2007b).

图3 扬子地块东部燕山期 (a-c)、天山北部地区石炭纪 (d-f) 和藏南冈底斯渐新世-中新世 (g-i) 埃达克质岩的SiO₂-MgO, SiO₂-Mg#和SiO₂-FeO*/MgO图解

Fig. 3 SiO₂-MgO, SiO₂-Mg# and SiO₂-FeO*/MgO diagrams for adakitic rocks of Yanshanian ages in the eastern Yangtze Block (a-c), Carboniferous ages in the northern Tianshan area (d-f) and Oligocene-Miocene ages in the Gangdese area (South Tibet) (g-i)

5 小结

1) “埃达克质岩”主要是指那些具有与俯冲洋壳熔融形成的“埃达克岩”类似主微量地球化学特征的岩浆岩。

2) 埃达克质岩有利成矿的背景主要包括岛弧、大陆板内伸展和活动大陆碰撞造山带环境。埃达克质岩的成矿潜力巨大。

3) 在岛弧和大陆板内伸展环境中, 熔体或超临界流体与地幔的相互作用可能是导致埃达克质

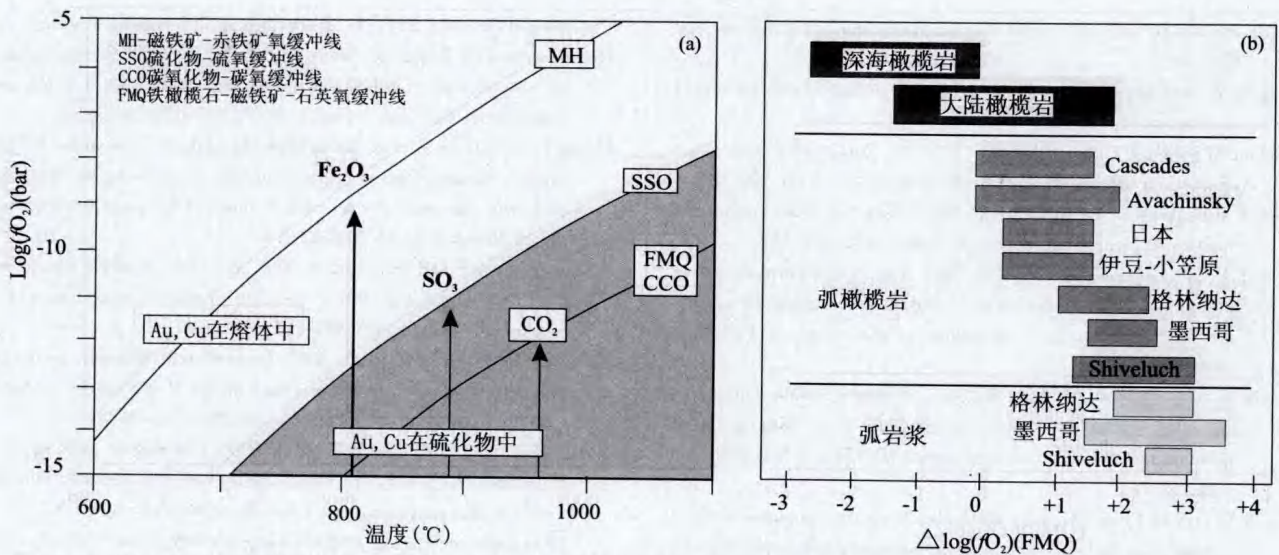


图4 $\log f_{O_2}$ -温度图解中的氧缓冲线 (Mungall, 2002) (a) 和橄榄岩、弧岩浆的氧逸度 (Bryant et al, 2007) (b)
 In Fig. a, only addition of Fe_2O_3 can bring mantle assemblages higher than SSO buffer. Undersaturation of magma in sulfide requires either very high degrees of partial melting or generation in unshaded portion of diagram. SSO-sulfide-sulfur oxide buffer, CCO-carbon dioxide-carbon oxide buffer, FMQ-fayalite-magnetite-quartz oxygen buffer; MH-magnetite-nematite oxygen buffer.

图4 $\log f_{O_2}$ -温度图解中的氧缓冲线 (Mungall, 2002) (a) 和橄榄岩、弧岩浆的氧逸度 (Bryant et al, 2007) (b)
 Fig. 4 Oxygen buffers in $\log f_{O_2}$ versus temperature space (a) and oxygen fugacity for peridotites and arc lavas (b)

岩有利于铜金等金属矿化的重要原因之一。

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